

## **Climate and statistics**

## Outline

### Anomalies Comparing climate models to data Downscaling and bias correction

## The fit of models

- Climate is the distribution of weather
- To reasonably estimate a distribution (from data or from models) need a relatively long stretch of data–WMO suggests at least 30 years

How well does the CMIP5 experiment used in the recent IPCC assessment work for describing annual global mean temperature?

## The greenhouse effect

Joseph Fourier (1768-1830) realized that Earth ought to be a lot cooler than it is. John Tyndall (1820-1893) found that water vapor and  $CO_2$  are greenhouse gases Svante Arrhenius (1859-1927) calculated how changes in  $CO_2$  can heat the planet







## A simple climate model



## Solution

 $T^{4} = \frac{1365 \times 0.7}{4 \times 0.61 \times 5.67} \times 10^{8}$ 

## Average earth temperature is T=288K (15°C)

One degree Celsius change in average earth temperature is obtained by changing

- solar constant by 1.4%
- Earth's albedo by 4.5%
- effective emissivity by 1.4%

## But in reality...

The solar constant is not constant

- The albedo changes with land use changes, ice melting and cloudiness
- The emissivity changes with greenhouse gas changes and cloudiness

Need to model the threedimensional (at least) atmosphere

But the atmosphere interacts with land surfaces...

...and with oceans!

## The climate engine I

If Earth did not rotate: tropics get higher solar radiation hot air rises, reducing surface pressure and increasing pressure higher up forces air towards poles lower surface pressure at poles makes air sink moves back towards tropics



## The climate engine II

- Since earth does rotate, air packets do not follow longitude lines (Coriolis effect)
- Speed of rotation highest at equator
- Winds travelling polewards get a bigger and bigger westerly speed (jet streams)
- Air becomes unstable
- Waves develop in the westerly flow (low pressure systems over Northern Europe)
- Mixes warm tropical air with cold polar air
- Net transport of heat polewards



# Modeling the atmosphere

#### Lagrangean approach



## Modeling, cont.



## The issue of gridding



Hurricanes Clouds Glaciers

## Comparing two distributions

Location shift:  $X + \delta \sim Y$ Location-scale:  $\alpha X + \delta \sim Y$ More general:  $X + \Delta(X) \sim Y$ Having m observations from Y and n observations from X we estimate  $\hat{\Delta}(x) = G_m^{-1}(F_n(x)) - x$ 



# Comparing global climate models to data



x Quantiles

### **Anomalies**

Comparison to "normal" Normal = 30 yr average Different baselines Helps for regional trends Really residuals So fit a model (trend + seasonal + covariates+variability)

## Global mean temperature





## **30-year distributions**

1930-1959

1930-1959



1970-1999 CCSM4 HadCM3 1970-1999 1970-1999 0.6 0.4 0.2 2 Frequenc§<sup>ISS data</sup> Data 0.2 , Ø Frequency 0.0 0.0 £0 € -0.2 4 0.2 110 0.2 0.6 0.4 06 -0.1 0.1 0.3 0.5  $\sim$ 50 CMIP5 Model Model m odels 0 0 0.8 -0.2 0.2 0.0 0.4 0.0 0.4 0.6 -0.4

CMIP5 models

GISS data

## Comparing location and spread



## **Effect of anomalies**

NCEI





End year

## **Dynamical downscaling**

#### Global models are very coarse Regional models are driven by boundary conditions given by global model runs



Jan/1/2000 0:00



## Swedish temperature minima

SMHI synoptic stations in south central Sweden, 1961-2008. SMHI regional model (open air & snow) Seasonal minima (d=1 DJF, d=2 MAM, d=3 JJA, d=4 SON).





### **Spatial models**

-m,(s) ~ GEV(μ,(s),σ(s),ξ) where  $\mu_{t}(s) = \beta_{0}(s) + \beta_{1}(s)(t - 1961) / 50$  $+\sum_{d=1}^{4}\beta_{d}(s)\mathbf{1}(d_{t}=d)$ d=2 $\beta_i(s) \sim GP(\mu_i, \sigma_i(1 - exp(-\theta_i d(s))))$  $\log \sigma(s) \sim GP(\mu, \sigma(1 - \exp(-\theta d(s))))$ Both for data and model output.

## RCM temperature minima





### Fit of GEV-distribution







## (Dis)agreement between RCM and data

Seasonal effects quite similar Similar spatial scale Similar shape parameter Temporal trend substantially lower in model output Data trend about 0.4-1°C /decade (lower than the annual model)

# Norwegian winter precip



## **Bias correction**

Need downscaled precipitation projections for adaptation plans Bias correction for downscaled

reanalysis

Apply to downscaling historical GCM

If works, apply to downscaling GCM projections

Correction more important for large quantiles than for entire distribution

## **Full quantile correction**

#### Applying the Doksum shift we get

$$Z_{it'}^{cal} = Z_{it'}^{H} + \hat{\Delta}_{i}(Z_{it'}^{H})$$
  
=  $Z_{it'}^{H} + \hat{G}_{i}^{-1}(\hat{F}_{i}(Z_{it'}^{H})) - Z_{it}^{H}$   
=  $\hat{G}_{i}^{-1}(\hat{F}_{i}(Z_{it'}^{H}))$ 

Rejections (fit to 80%, tested on 20%):

Raw 77%

**Corrected 18%** 

**Corrected GCM 79%** 

## Single quantile correction



## $log(Y^{q}) = \alpha + diag(log(X^{q}))\beta + \varepsilon$ $\beta \sim GP(0, \Sigma(\nu, \kappa, \phi))$

Replace with smoothed version

## **Bias-correction of 95<sup>th</sup> precipitation percentile**

